

Impact of the ocean-atmosphere coupling on high-resolution future projections for the Mediterranean sea and surrounding climate from the Med-CORDEX ensemble

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Med-CORDEX

International initiative that aims at developing **fully coupled high resolution Regional Climate Models (RCMs) for the Mediterranean basin**, as part of the global CORDEX initiative.

Institutions

- CMCC
- CNRM
- ENEA
- GERICS-AWI
- GUF
- LMD
- U. Belgrade

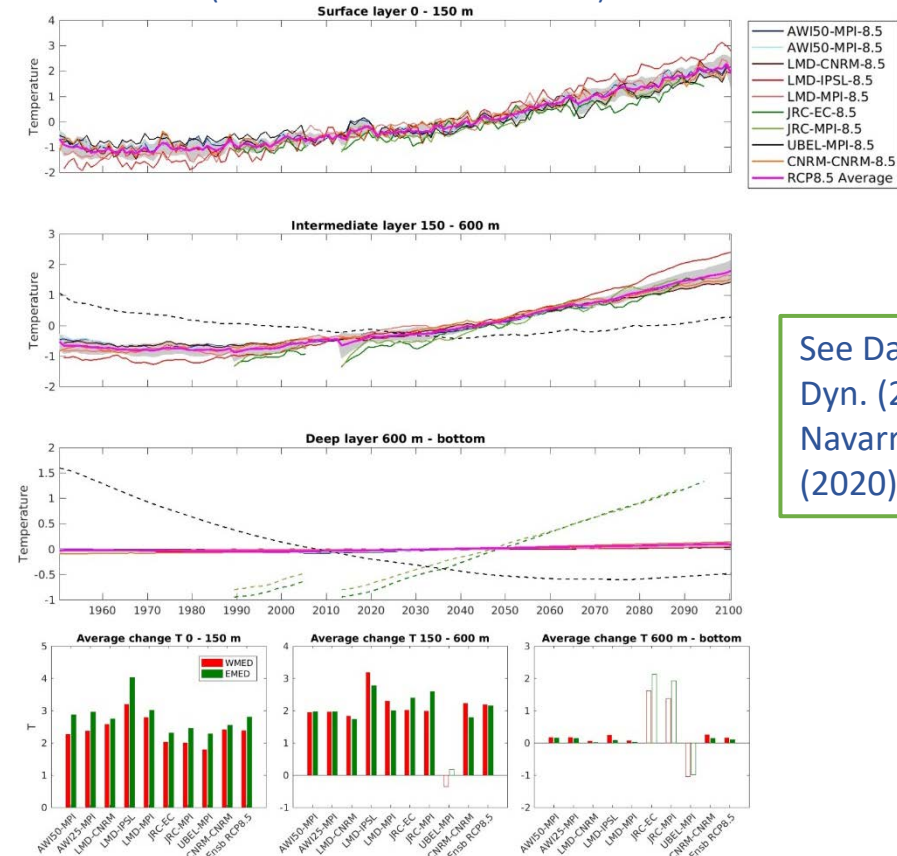
Climate change studies
in the Mediterranean
basin



Baseline atm – ocean
coupled simulations

26 historical and
multi-scenario
simulations

Temperature anomalies of the water column from 1950 to 2100 (°C)
(from Soto-Navarro et al 2020)



See Darmaraki et al., Clim. Dyn. (2019) and Soto-Navarro et al., Clim. Dyn. (2020) for details

Objectives:

- Is the climate change response consistent in all the models?
- Is there a significant impact of the higher resolution in the RCMs?
- Is there a significant impact of the ocean-atmosphere coupling?

In this presentation: 14 simulations (7 historical, 7 scenario runs)

Institution	RCM	ARCM	GCM	Scenario	Short Name
CNRM	CNRM-RCSM4		CNRM-CM5	RCP 8.5	CNRM-CM5
	CNRM-RCSM6	CNRM-ALADIN63	CNRM-ESM2-1	SSP 5-85	CNRM-ESM2
GERICS-AWI	GERICS-AWI-ROM25	REMO25	MPI-ESM-LR	RCP 8.5	AWI-25-MPI
	GERICS-AWI-ROM50	REMO50	MPI-ESM-LR	RCP 8.5	AWI-50-MPI
ENEA	ENEA-PROTHEUS		CNRM-CM5	RCP 4.5	ENEA-CNRM
LMD	LMD-LMDZNEMOMED8		IPSL-CM5A-MR	RCP 8.5	LMD-IPSL
U. Belgrade	EBU-POM2	EBU	MPI-ESM-LR	RCP 8.5	UBEL-MPI

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In this presentation: 14 simulations (7 historical, 7 scenario runs)

- Analysis of the climate change signal of SST and atmospheric variables at the surface level

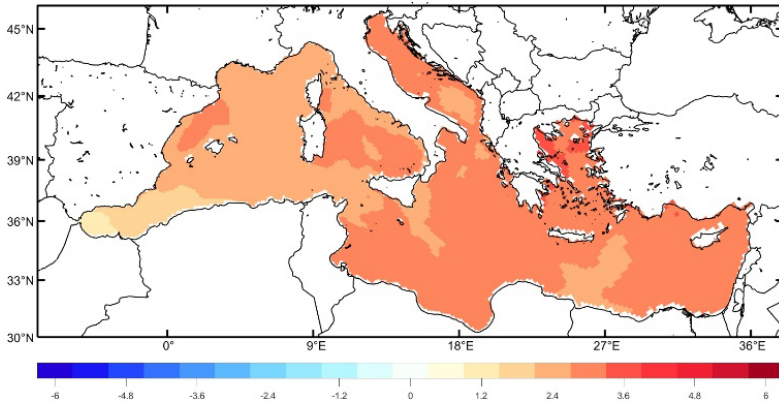
Climate Change (CC) signal computed as the difference between the averages of the last 30 years of the projection (2070-2100) and the last 30 years of the historical period (1976-2005).

$$\text{CC signal} = \text{average}(2070-2100) - \text{average}(1976-2005)$$

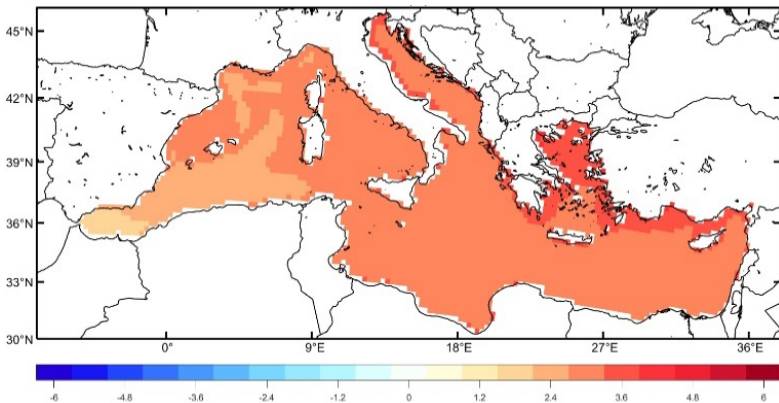
Models CC response

SST and Air temperature increase

AWI-25-MPI SST CC signal (°C)



AWI-25-MPI Air T CC signal (°C)



- All the simulations show a warming of the sea surface between 2.5 and 4 °C on average (1.2-1.5 °C for RCP 4.5)
- The Air T average increase is around 30% higher than for the SST

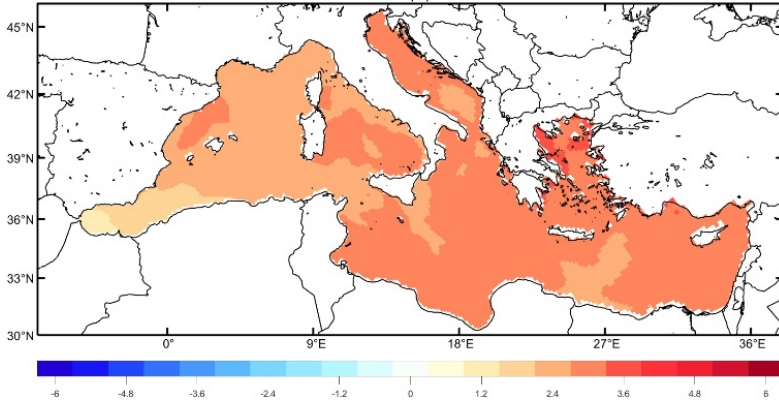
Models CC response

SST and Air temperature increase

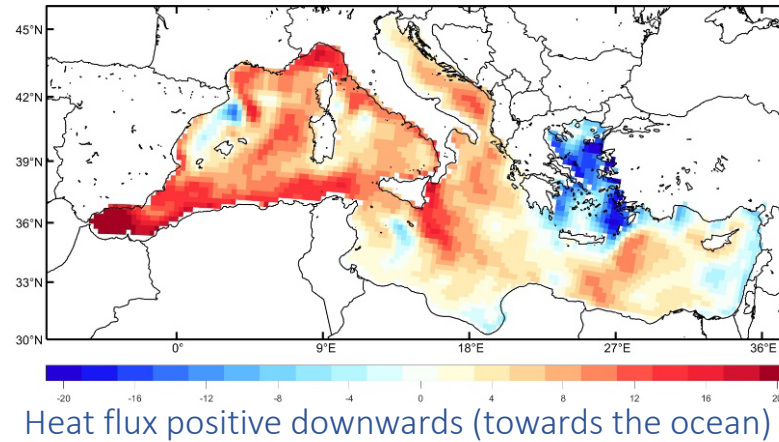


Decrease of the net heat losses
towards the atmosphere

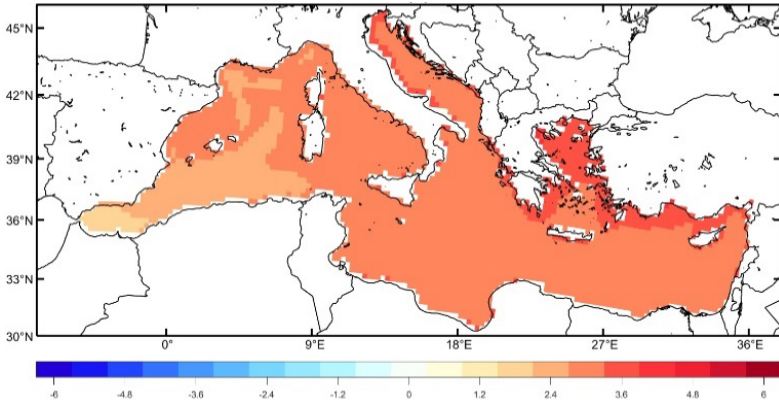
AWI-25-MPI SST CC signal (°C)



AWI-25-MPI Net Surface Heat Flux CC signal (W/m²)



AWI-25-MPI Air T CC signal (°C)



- Decrease in the net heat loss (average 0.2 – 4.3 W/m²) with high spatial variability → the atmosphere is cooling less the sea and even starting to warm it for some models

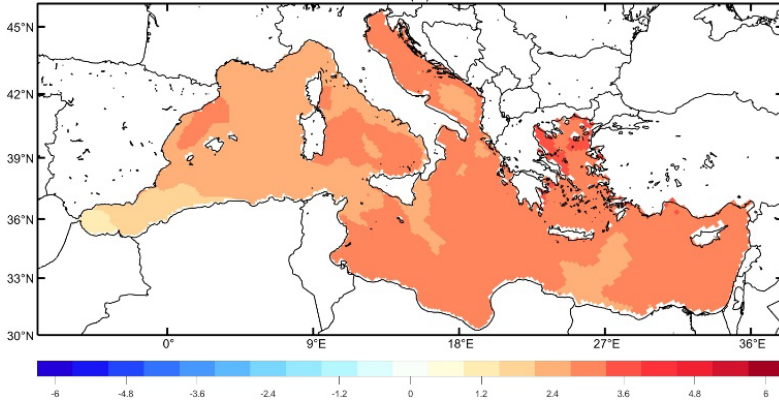
Models CC response

SST and Air temperature increase

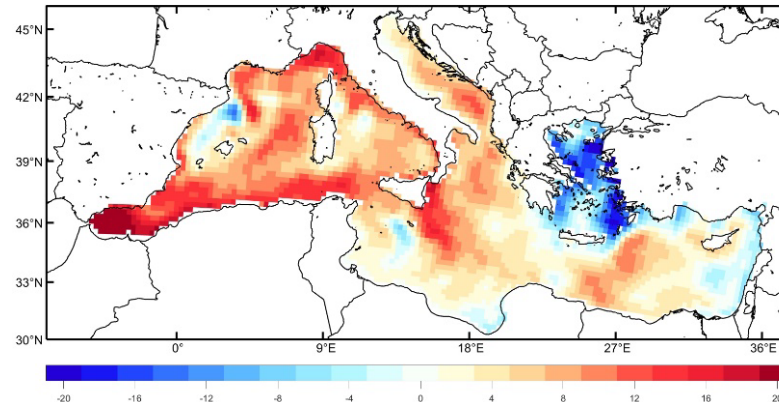
Decrease of the net heat losses
towards the atmosphere

Decrease of Precipitation

AWI-25-MPI SST CC signal (°C)

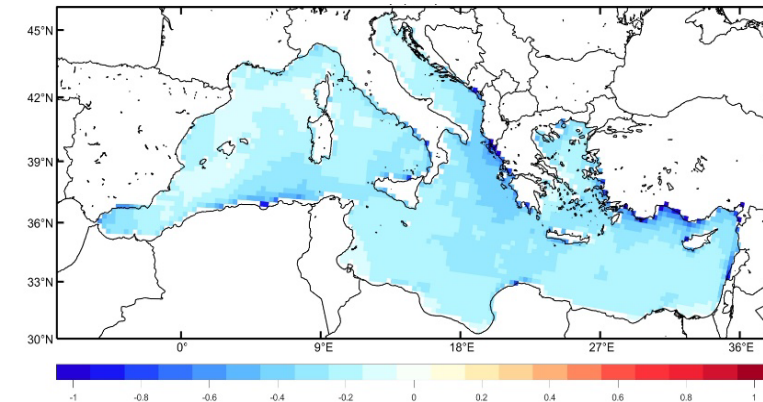


AWI-25-MPI Net Surface Heat Flux CC signal (W/m²)

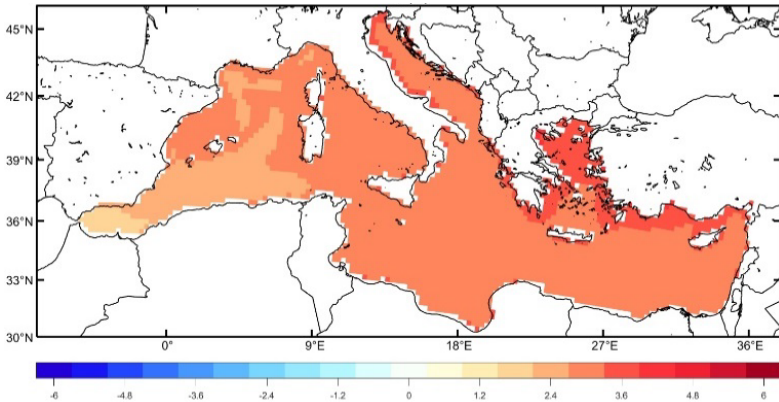


Heat flux positive downwards (towards the ocean)

AWI-25-MPI Precipitation CC signal (mm/d)



AWI-25-MPI Air T CC signal (°C)



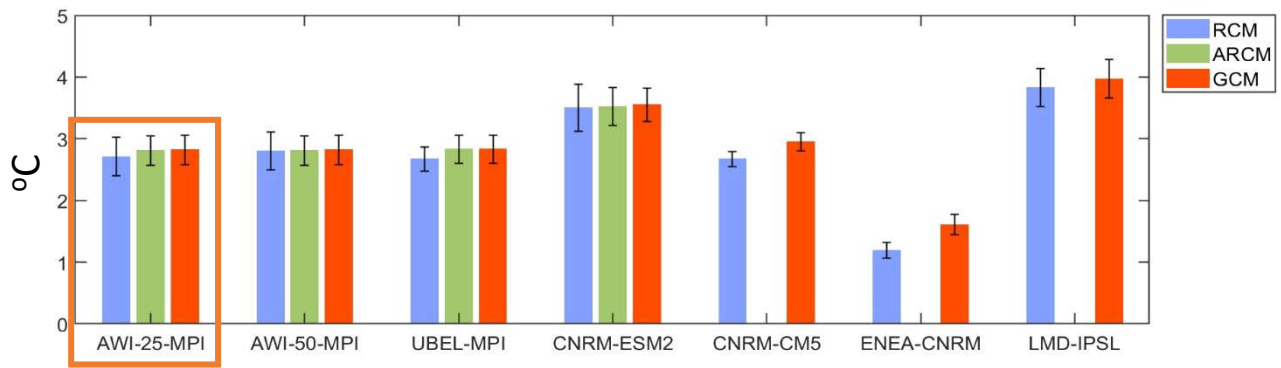
➤ Precipitation decreases in all RCMs
(average -0.1 – -0.4 mm/d)

➤ Decrease in the net heat loss (average 0.2 – 4.3 W/m²) with high spatial variability → the atmosphere is cooling less the sea and even starting to warm it for some models

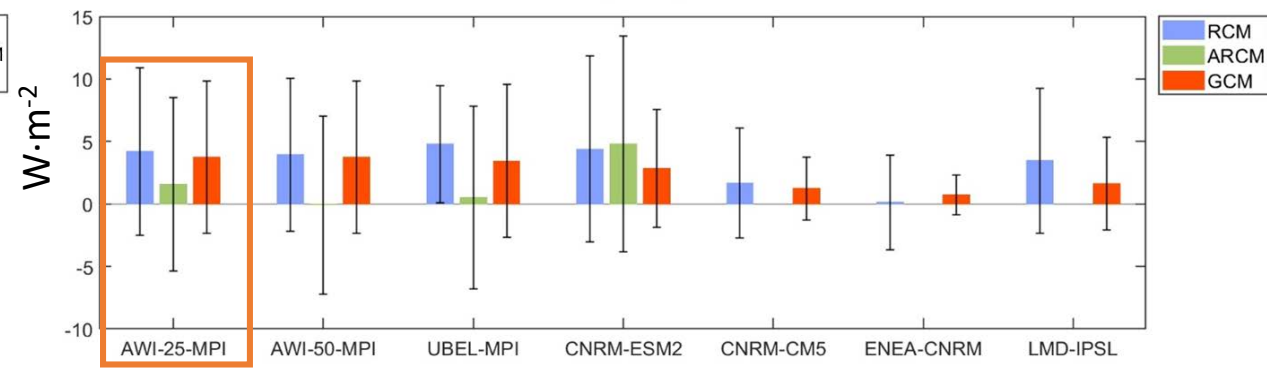
Models CC response

Consistency among all the models

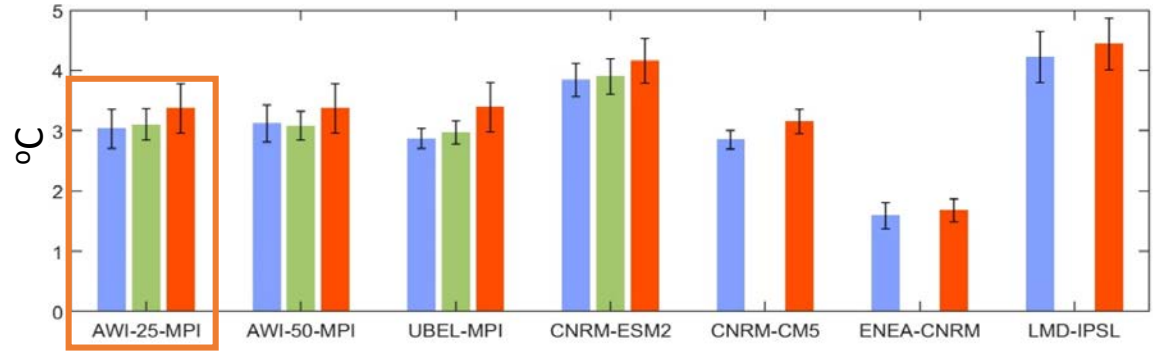
SST CC signal



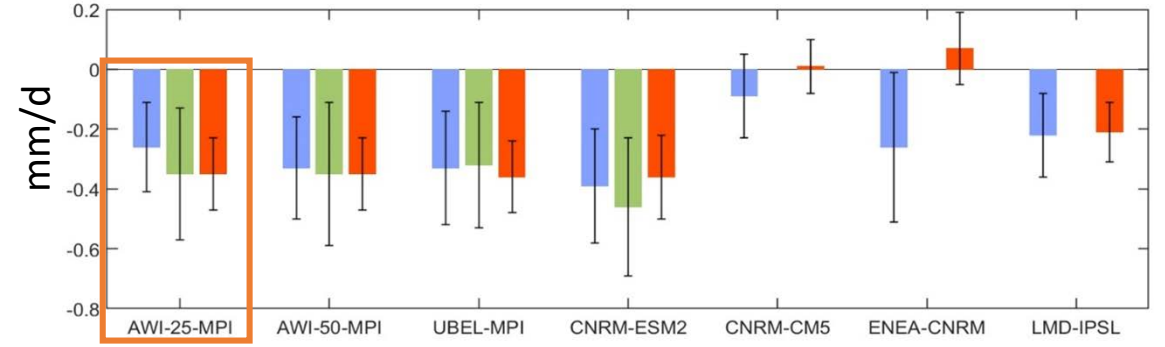
Net Surface Heat Flux CC signal



Air T CC signal



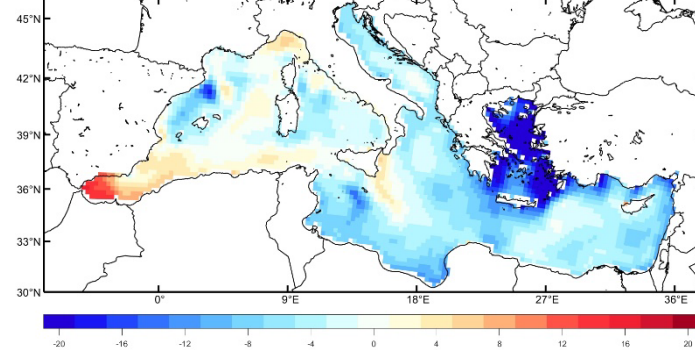
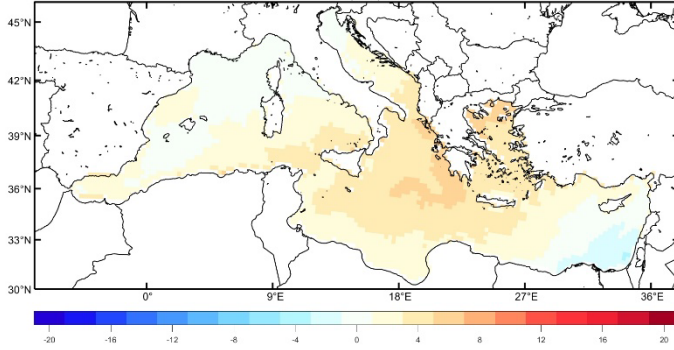
Precipitation CC signal



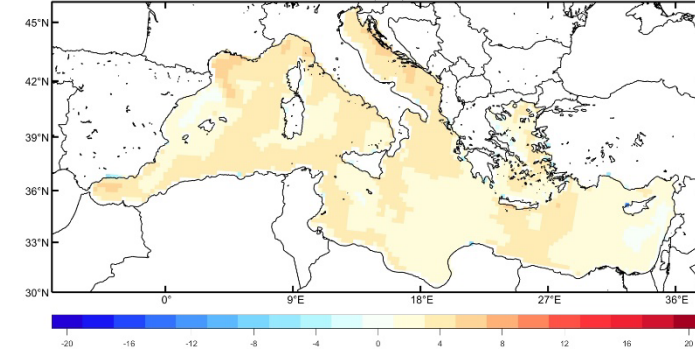
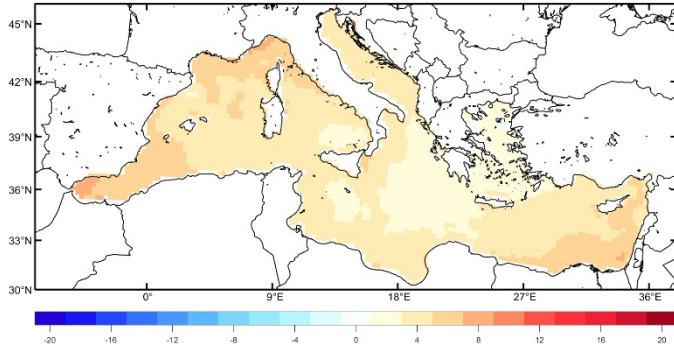
Models CC response

Surface heat flux components

AWI-25-MPI Net Shortwave Radiation CC signal (W/m²) AWI-25-MPI Latent Heat Flux CC signal (W/m²)



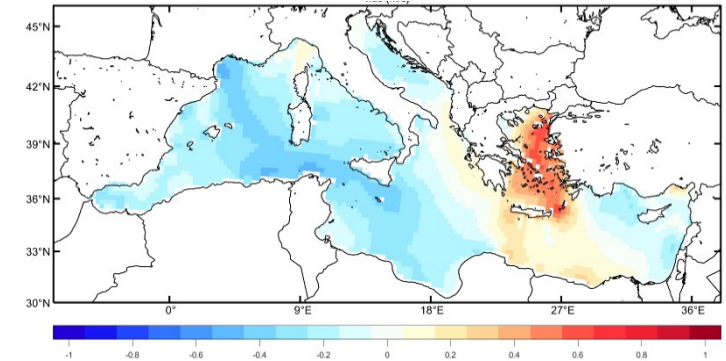
AWI-25-MPI Net Longwave Radiation CC signal (W/m²) AWI-25-MPI Sensible Heat Flux CC signal (W/m²)



Heat flux positive downwards (towards the ocean)

- Increase of the net **shortwave** radiation → reduction in the cloud cover and aerosol concentration
- Increase of the net **longwave** radiation → atmospheric warming due to GHG (among other factors)
- Increase of the **sensible** heat flux → due to the increase in the difference between the Air and Sea Temp
- Decrease of the **latent** heat flux, meaning an **increase** in the heat loss due to **evaporation** → the only negative term (tend compensate). **Modulated by wind**

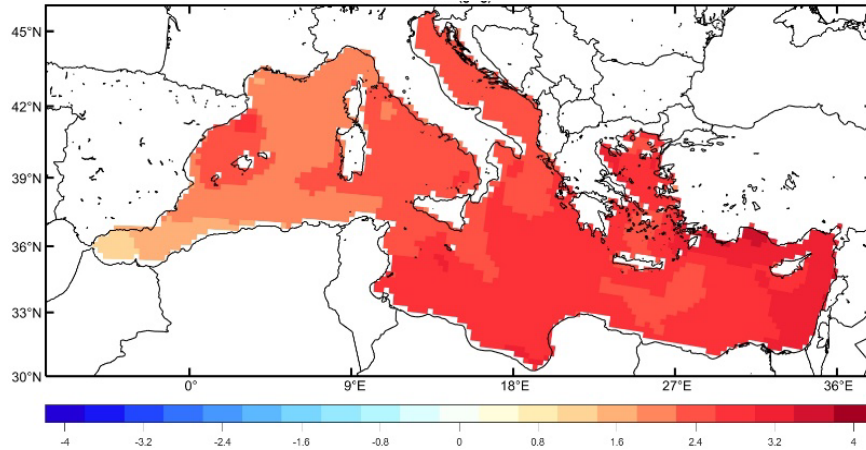
AWI-25-MPI Wind Intensity CC signal (m/s)



Models CC response

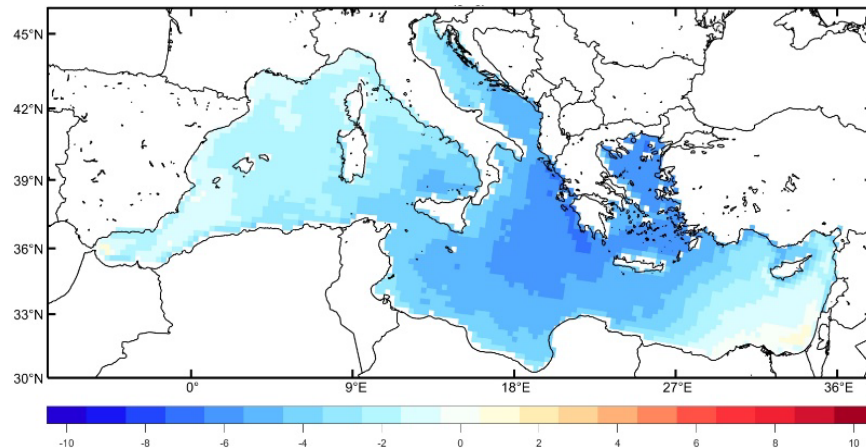
Humidity and clouds cover

AWI-25-MPI Specific Humidity CC signal (g/kg)



- Increase of the specific humidity due to the increase of the evaporation and atmospheric warming (Clausius-Clapeyron)
- General decrease in the cloud cover → Med climate is less prone to convection in the future

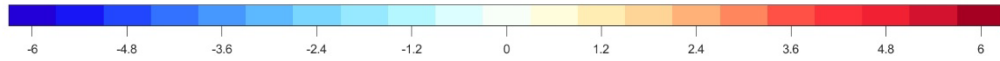
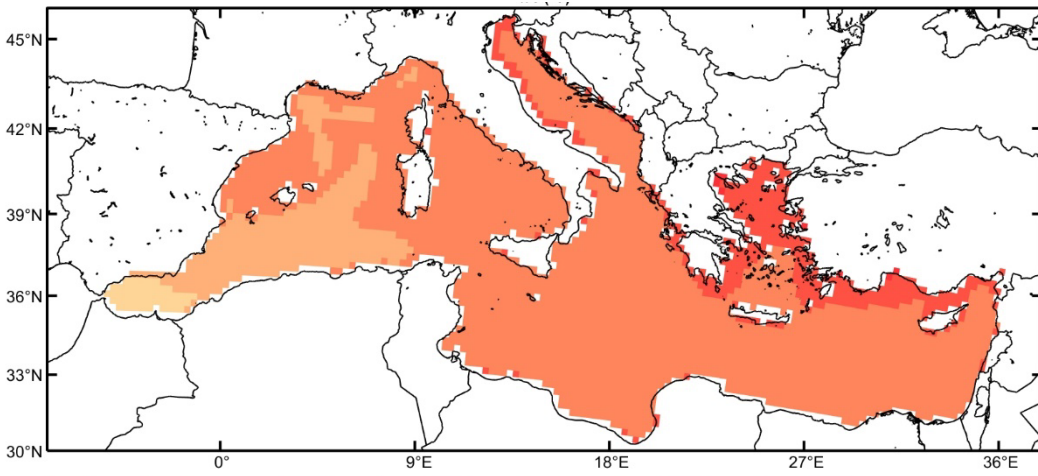
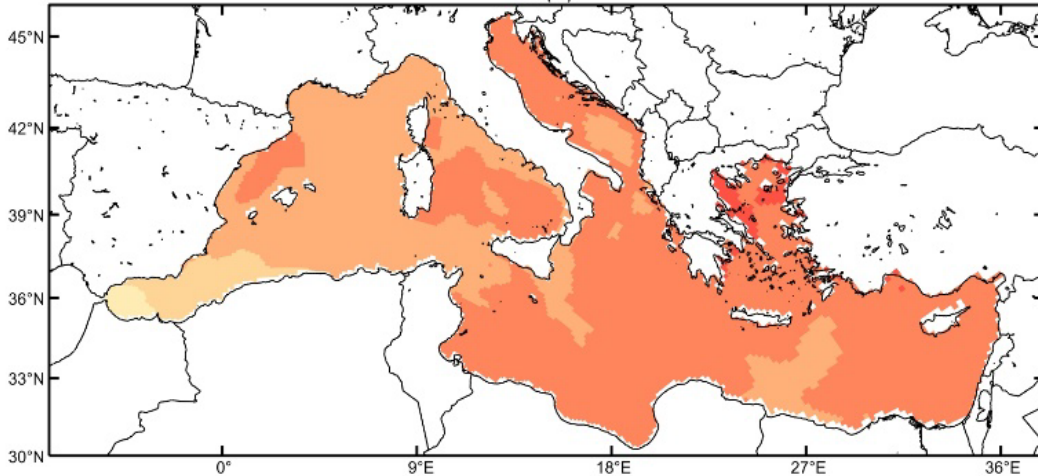
AWI-25-MPI Clouds cover CC signal (%)



Impact of the resolution

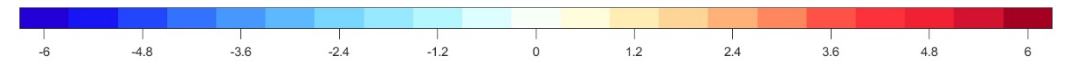
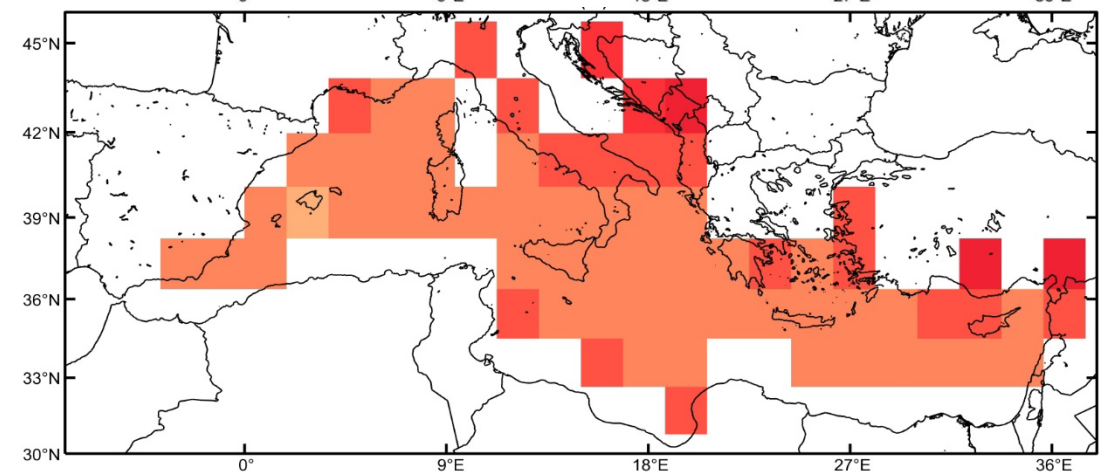
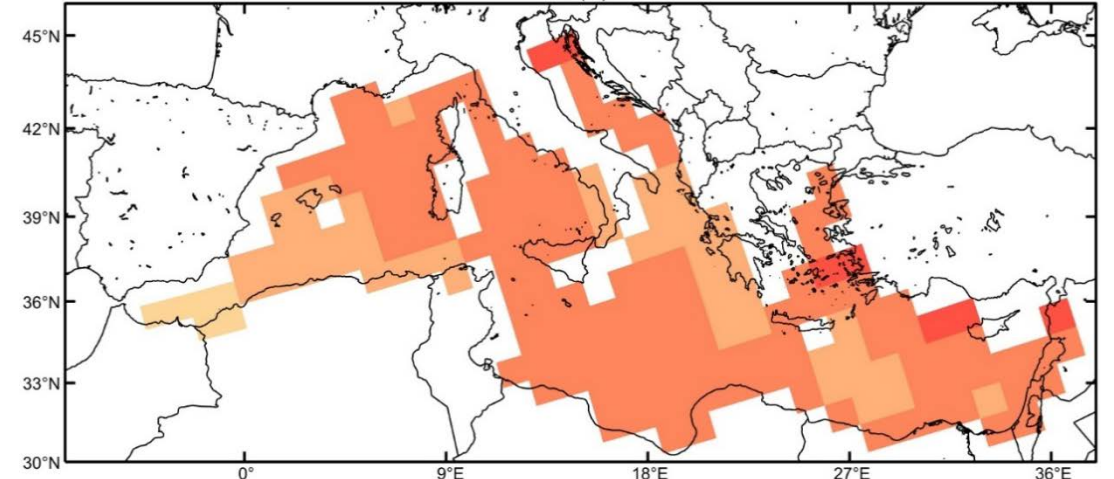
RCM

AWI-ROM25 SST & Air T CC signal (°C)



GCM

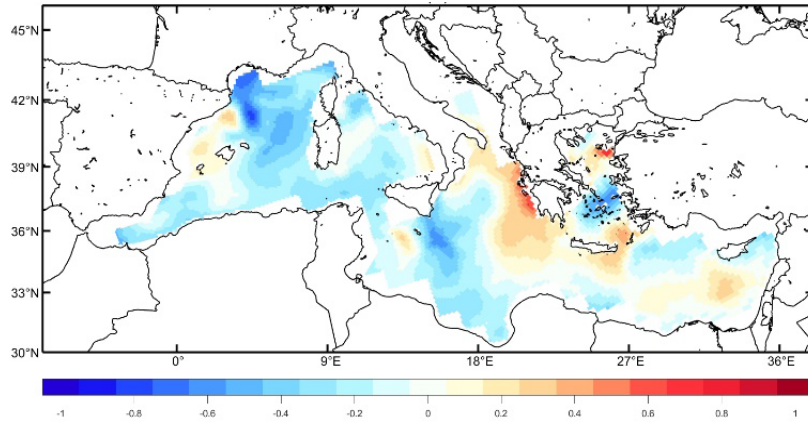
MPI-ESM-LR SST & Air T CC signal (°C)



Impact of the resolution

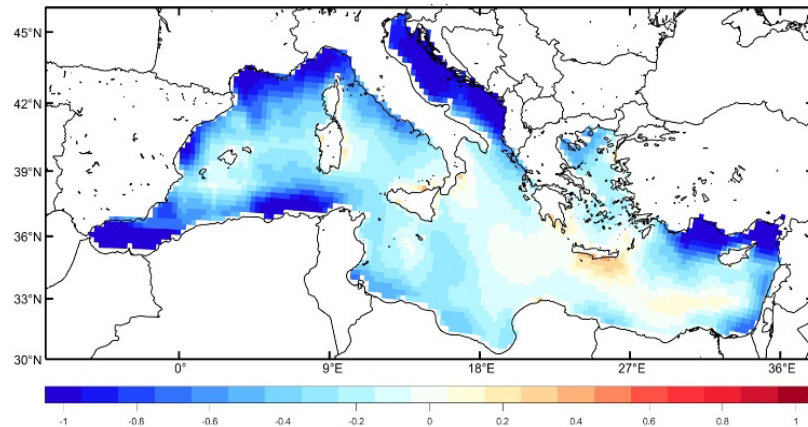
Differences between the RCMs and GCMs CC signals

AWI-ROM25 vs MPI-ESM-LR SST dCC signals (°C)



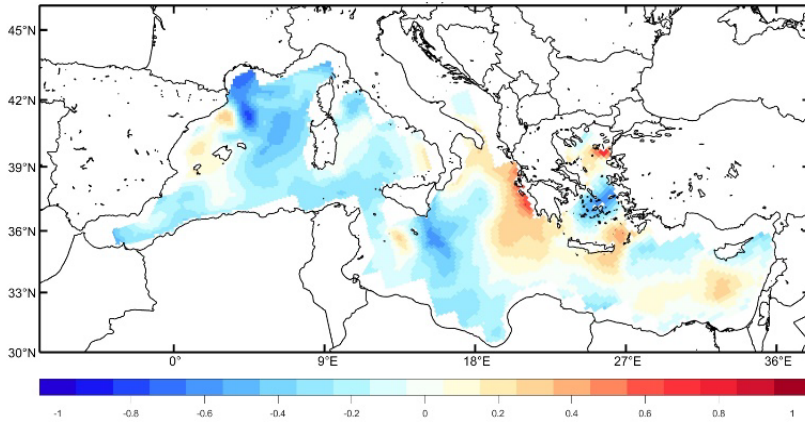
- The average SST CC signal is slightly stronger in the GCMs in general
- Significant changes in the spatial structures
- The air T signal is clearly stronger in the GCMs

AWI-ROM25 vs MPI-ESM-LR Air T dCC signals (°C)

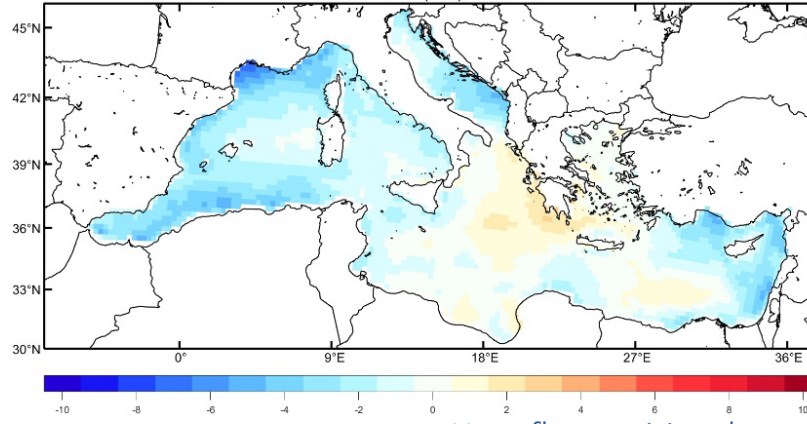


Differences between the RCMs and GCMs CC signals

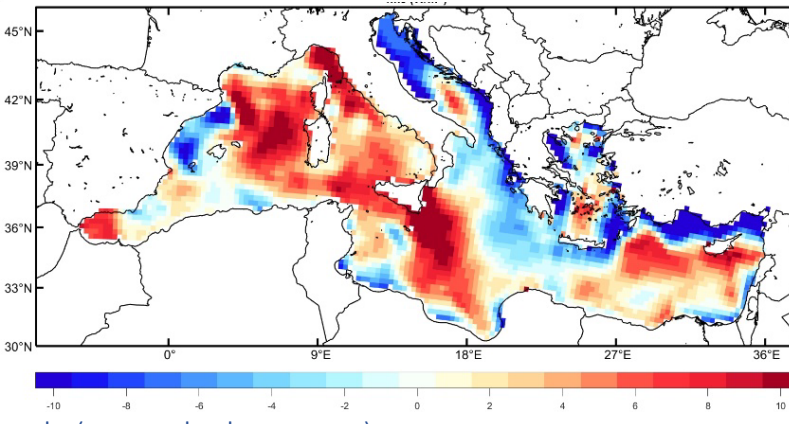
AWI-ROM25 vs MPI-ESM-LR SST dCC signals (°C)



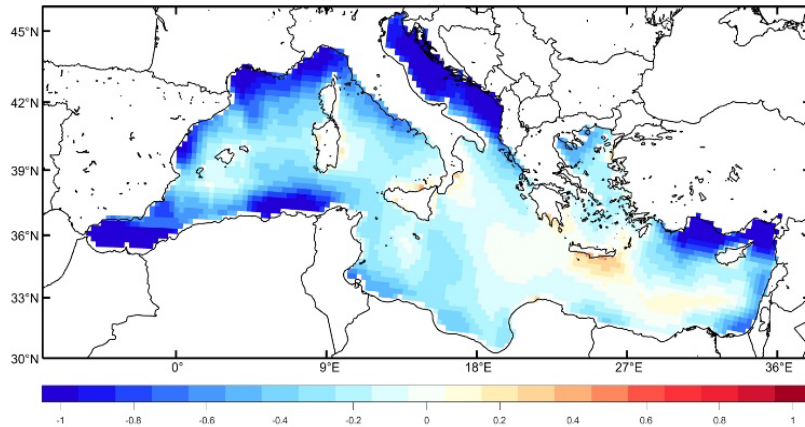
AWI-ROM25 vs MPI-ESM-LR NSWR dCC signals (W/m²)



AWI-ROM25 vs MPI-ESM-LR HFLS dCC signals (w/m²)



AWI-ROM25 vs MPI-ESM-LR Air T dCC signals (°C)



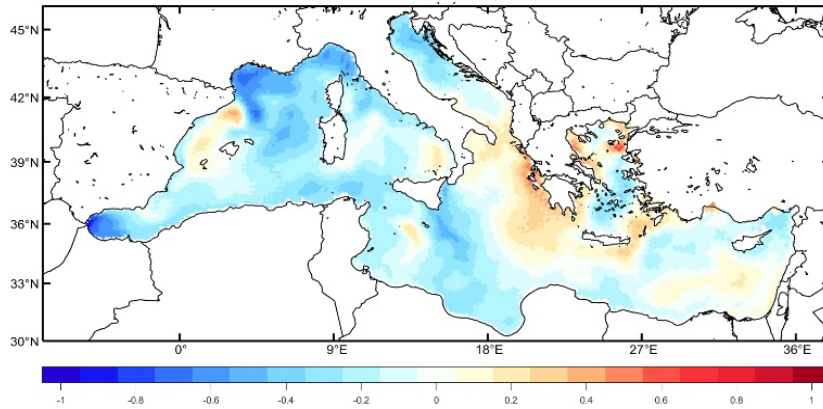
Heat flux positive downwards (towards the ocean)

- RCMs signals weaker than GCMs for shortwave rad → higher reduction of the cloud cover in the GCMs and not inclusion of aerosols in all RCMs
- RCMs signal weaker than GCMs for latent heat flux → more evaporation in the GCMs due to higher SST increase
- RCMs sensible heat flux signals are stronger than for the GCMs in 5 of the 7 simulations (not shown) → stronger gradient between SST and Air-T in RCMs
- No consistent difference between RCMs and GCMs in Precipitation

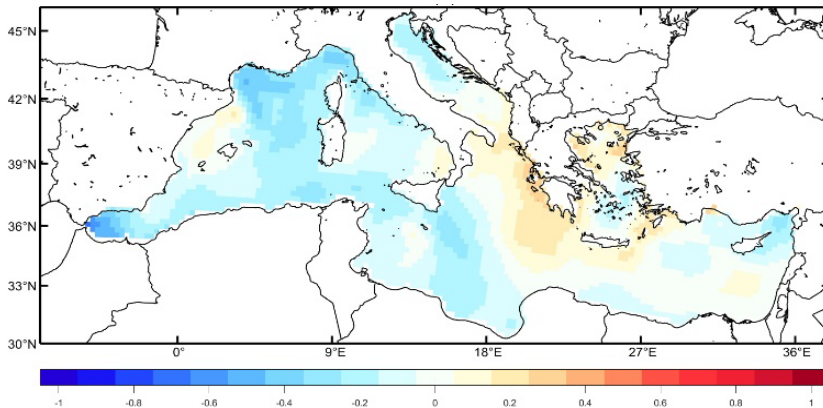
Impact of the coupling

Differences between the RCMs and ARCMs CC signals

AWI-ROM25 vs REMO25 SST dCC signals (°C)



AWI-ROM25 vs REMO25 Air T dCC signals (°C)

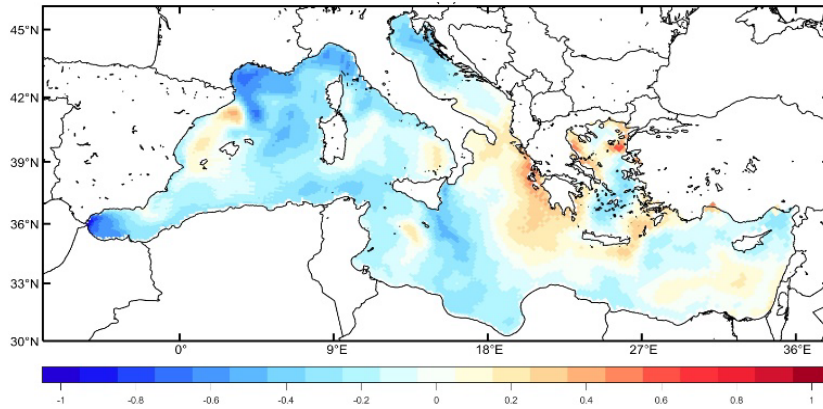


- The average SST CC differences are very close to the differences with the GCMs, as expected because the ARCMs use GCMs as boundary layer
- Significant changes in the spatial structures
- The air T signal differences are not as pronounced as with the GCMs, but still significant for the spatial structures

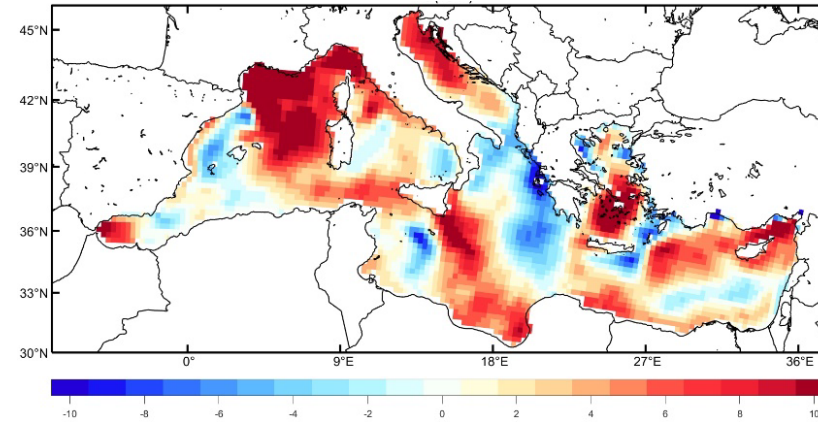
Impact of the coupling

Differences between the RCMs and ARCMs CC signals

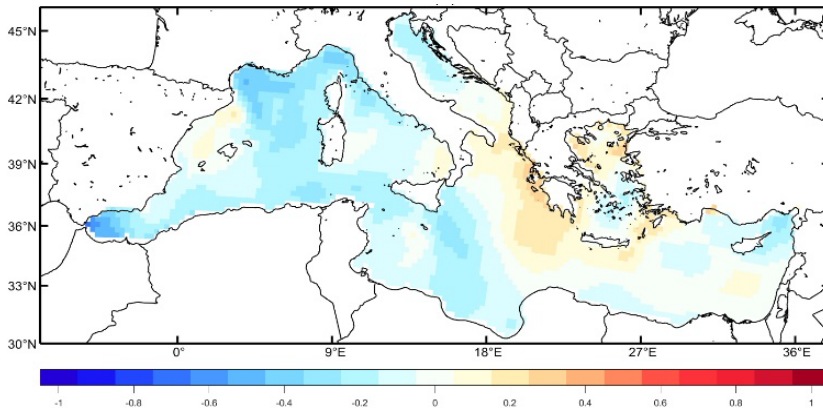
AWI-ROM25 vs REMO25 SST dCC signals (°C)



AWI-ROM25 vs REMO25 HFLS dCC signals (W/m²)



AWI-ROM25 vs REMO25 Air T dCC signals (°C)



- ARCMs show stronger latent HF signals than both RCMs and GCMs
→ more evaporation to compensate the higher SST increase from the GCMs boundary condition at the sea surface
- Also higher humidity increase in the ARCMs (not shown) → SST – Air -T gradient and latent heat flux differences
- No consistent difference between RCMs and ARCMs in Precipitation

Summary

- There is consistency in the CC response of all the simulations for the variables analyzed.
- The warming of the sea surface and the air results in a reduction of the net heat loss by the sea. The only component of the surface net heat flux that tends to counter this effect is the latent heat flux (increase of the ocean heat loss by evaporation).
- Despite the evaporation and humidity increase, the average cloud cover and precipitation decrease over the Mediterranean.
- Similar general behavior over land, but with much larger spatial variability.
- The main differences between RCMs and GCMs CC response are the SST and Air T signals, which in turn condition the ocean-atmosphere net heat flux.
- The RCMs dump this difference by the ocean-atmosphere interaction while in the ARCMs increase the latent heat flux losses (more evaporation) to compensate the extra sea surface warming from the GCMs boundary condition.
- Therefore, there is a significant impact of both the high resolution and the coupling in the models response to the climate change forcing.



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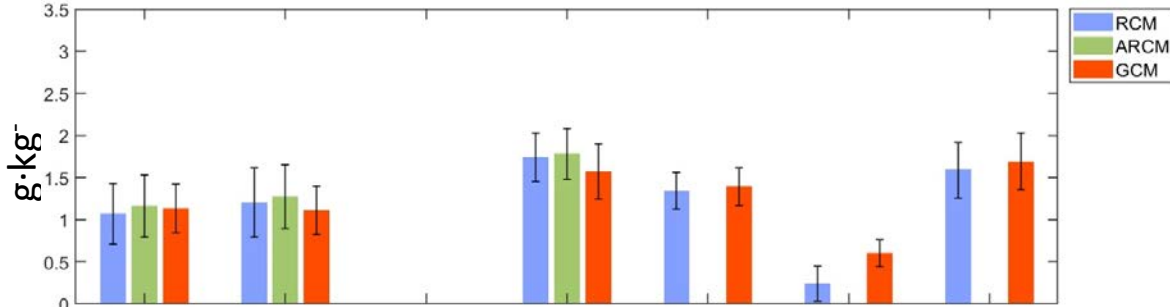


Extra slides

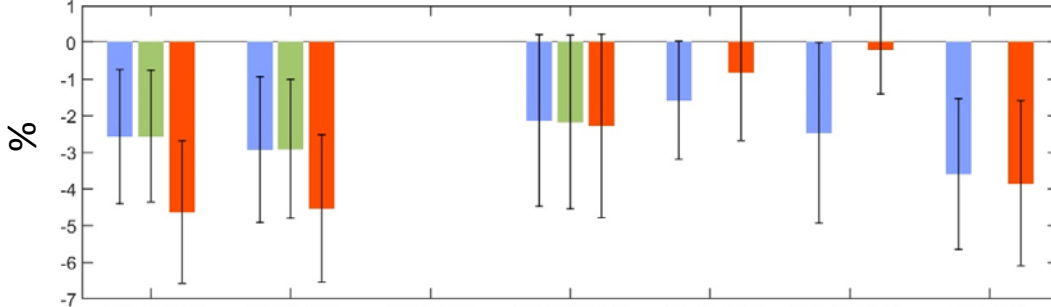
Impact over land

Spatial averages of the CC signal over land

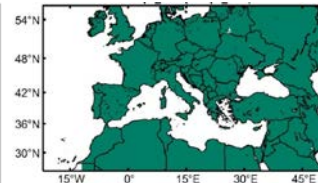
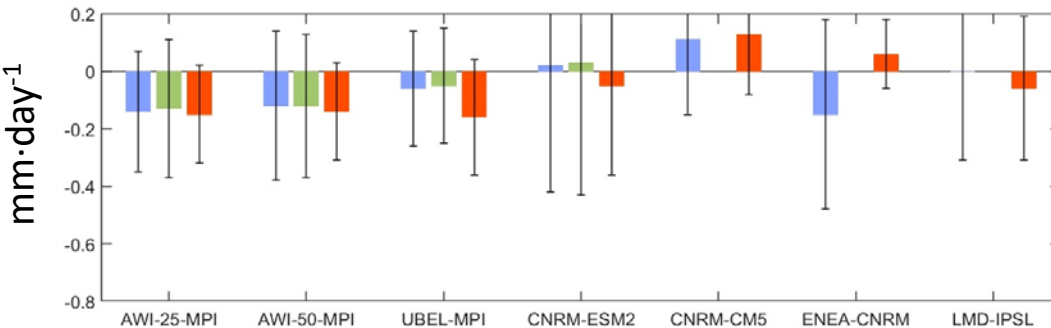
Specific humidity CC signal



Cloud cover CC signal



Precipitation CC signal



- As for the Med, general increase of humidity over land, but of lower magnitude. Higher in the RCMs and even more in the ARMs because of the HFSL differences
- Again, general decrease in the clouds cover, lower than over sea.
- Over land 3 of the 4 GCMs show a stronger decrease of the clouds cover. The differences between RCMs and ARCMs are no significant.
- The precipitation behavior is not as clear over land than over sea.
- All models show a very small average signal, but very high spatial variability.
- Clear diff between GCM response and RCM/ARCM.
- RCM seem to be drier than ARCM → change in land-sea contrast

The land region is too large and affected by different and complex processes on each region. More accurate results will be obtained when analyzing the land-sea interaction over specific land regions instead of the whole domain.